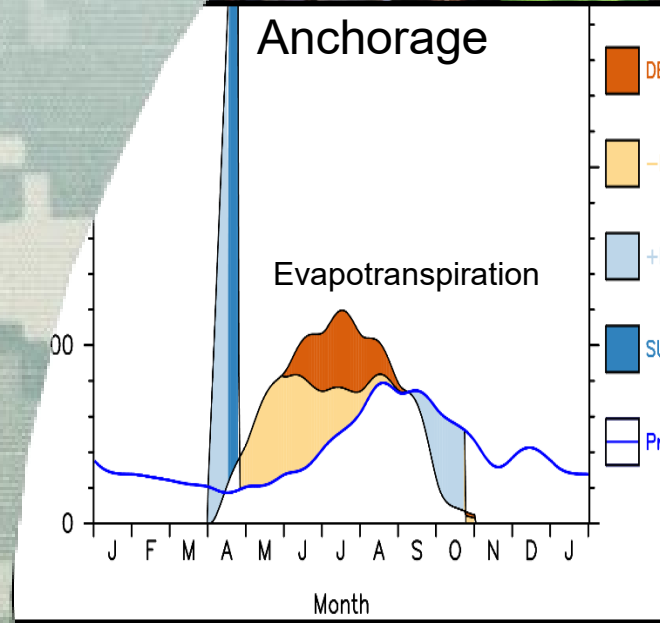


Alaska Region – REG IV Wetland Delineation

Soils Background



Objectives

The purpose of this presentation is to:

- ▶ discuss soil properties related to hydric soil development
- ▶ describe what information is required to be recorded on the soils data form
- ▶ provide guidance on describing these features.

This includes identifying soil layers, soil color, soil matrix, soil texture (organic vs mineral), redox features, determining percent of redox, where to begin observations, and understanding the concepts of reduced matrix and depleted matrix.

Definition of a Hydric Soil

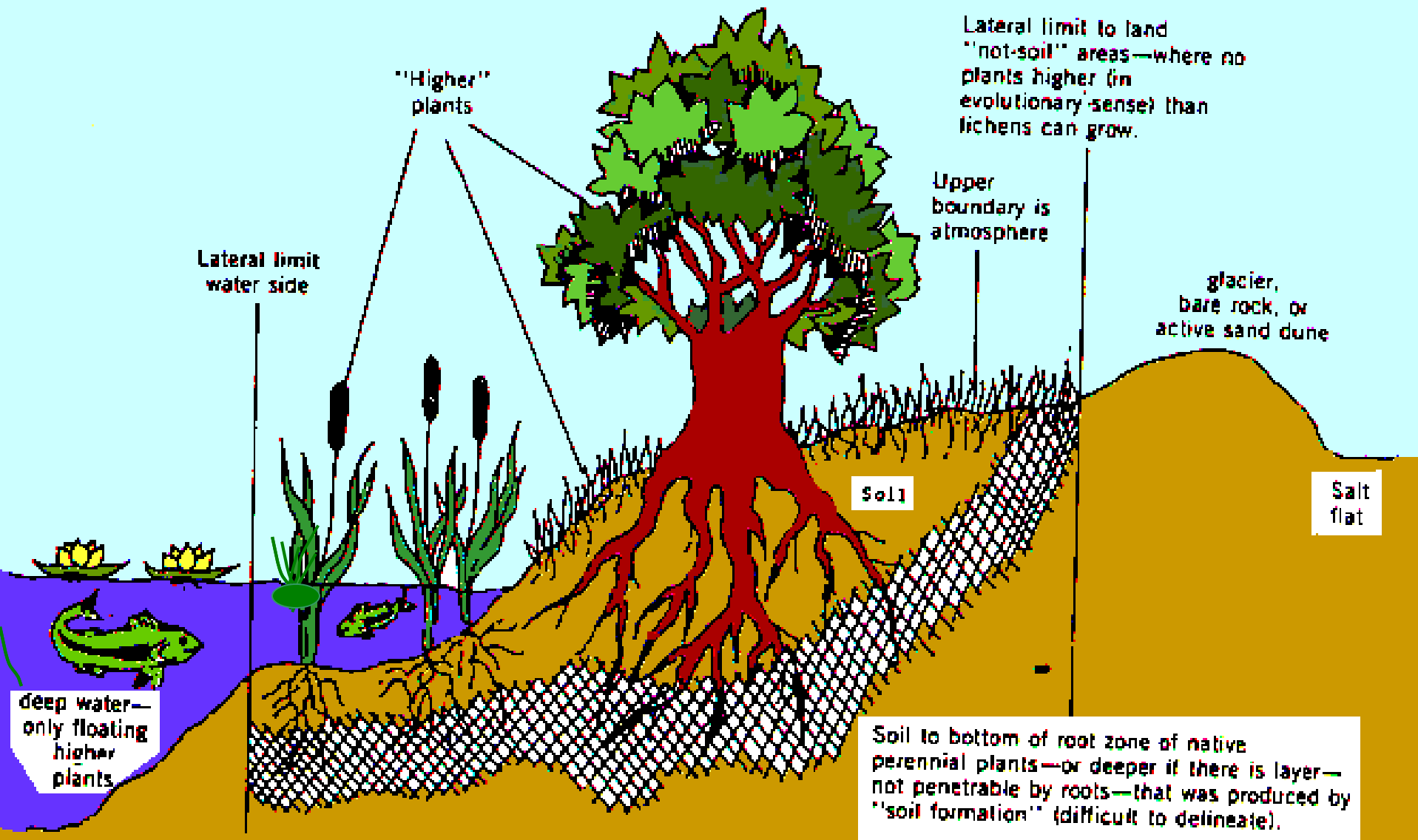
A hydric soil is a soil that formed under conditions of saturation, flooding, or ponding long enough during the growing season to develop anaerobic conditions in the upper part.



What is Soil?

Natural body that occurs on the land surface, occupies space, and is characterized by one or both of the following:

- Horizons or layers
- The ability to support rooted plants in a natural environment
 - ▶ Upper limit is air or shallow (>2.5 m) water
 - ▶ Lower limit is either bedrock or the limit of biological activity
 - ▶ Lower limit for classification set at an arbitrary 2 m



- Badlands
- Beaches
- Chutes
- Cinder land
- Dams
- Dumps
- Dune land
- Glaciers
- Gullied land
- Lava flows
- Mined land
- Oil-waste land
- Pits
- Playas
- Riverwash
- Rock outcrop
- Rubble lands
- Slickens
- Urban land
- Water

Non-soil



Defined in section 627 of the National Soil Survey Handbook

Factors That Influence Soil Development

- **Parent material:** soil derived from weathered sandstone is likely to have different drainage properties compared to clayey lacustrine sediments formed in glacial lakes.
- **Topography:** convex vs concave landscapes will drain and collect water respectively. Long term presence of water may result in formation of hydric soils.
- **Organisms:** Bioturbation from earthworms, animal burrows, and crawfish can influence development of macropores and soil structure.
- **Climate:** different climates allow for varying influences of weather on soil formation (a few months in Alaska vs year-round growing season in Florida).
- **Time:** Older soils tend to be thicker and have more layers or horizons than younger soils.

Parent Material

The material from which soil is formed

*Derived from physical and
chemical weathering of rocks
Produced by geologic or biologic
process*

- ▶ *Glacial materials*
- ▶ *Volcanic flows and ejecta*
- ▶ *Organic matter*
- ▶ *Tidal Sediments*

- ▶ *In place – residuum*
- ▶ *Moved by gravity - colluvium*
- ▶ *Moved by wind – loess*
- ▶ *Moved by water - alluvium*



Residuum/Colluvium



Glacial Materials

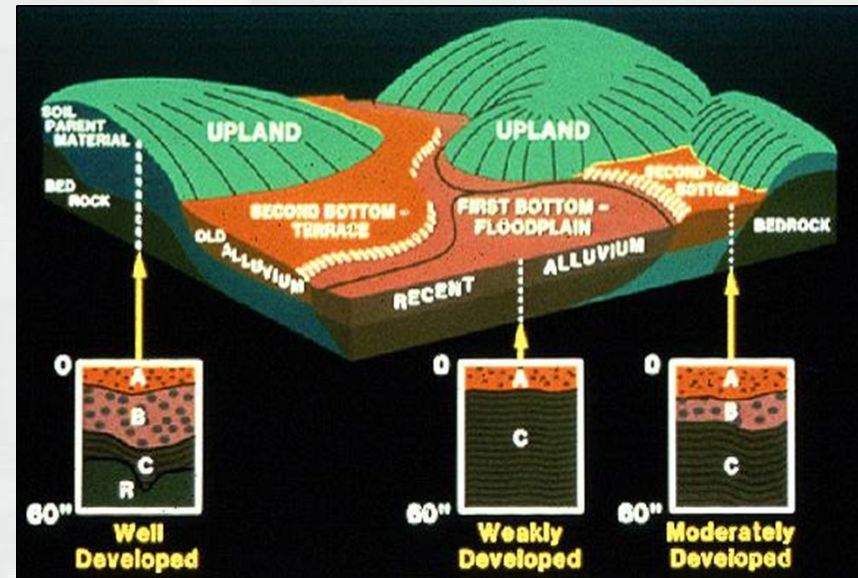
Topography

■ Landscape Position

- ▶ *Where are you?*
 - **Hill**
 - **Stream Terrace**

■ Landform

- ▶ **Glacial** – esker, kame, drumlin, moraine, outwash plain
- ▶ **Residual** – Bedrock cored Hill, Mountain
- ▶ **Wind** – dunes, loess hill
- ▶ **Water** – alluvial fan, floodplain, stream terrace, tidal flat
- ▶ **Permafrost** – thermokarst ponds, ice wedge polygons
- ▶ **Volcanic** – cinder cone, lava flow, volcanic cone



■ Shape

- ▶ **Concave**
- ▶ **Convex**
- ▶ **Plane**

■ Slope

■ Aspect

Effects of Landscape Position

- Critically influences water flow and soil formation
- Most wetlands are on some sort of concave surface
- Aspect: North vs. South

Hill Slope Elements and Curvature

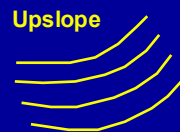
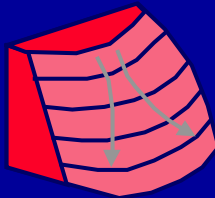
After Pennock et al., 1987

Slope

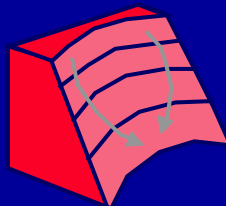
Block

Contour

Divergent



Convergent



Overland and Throughflow: Convergent landscapes

Potential hydric soil zone

Runoff

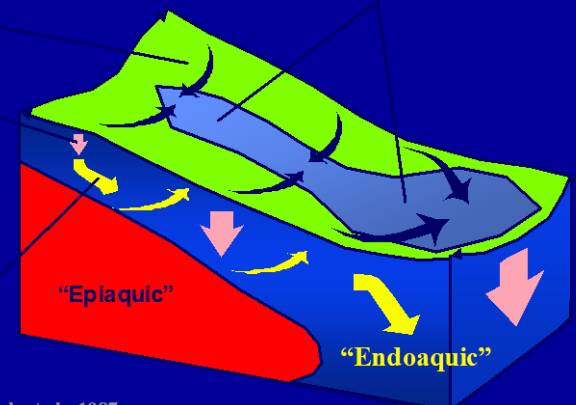
Infiltration
Percolation

Throughflow

"Epiaquic"

"Endoaquic"

Modified from Pennock et al., 1987



Organisms

■ Plants

- ▶ Trees
- ▶ Shrubs
- ▶ Herbs
- ▶ Mosses
- ▶ Lichens



Lichens on Rocks

■ Microorganisms

- ▶ Fungi
- ▶ Bacteria

■ Animals/Humans

- ▶ Excavations
- ▶ Organic Matter Decomposition
- ▶ Organic Matter Additions



Squirrel Midden

Climate

Climate is the major factor determining the kind of plant and animal life on and in the soil

■ Temperature

- ▶ Freezing and thawing
- ▶ Rate of chemical activity



■ Precipitation

- ▶ Wetting and drying
- ▶ Amount of chemical activity

Temperature and Precipitation fluctuations control the rate of physical and chemical weathering and the break down of parent materials in the soil



Time

Required for horizon formation

- Horizon formation is a slow process
- Long-term changes happen over hundreds, thousands, or millions of years
- Short-term changes can be caused by human activity and fire

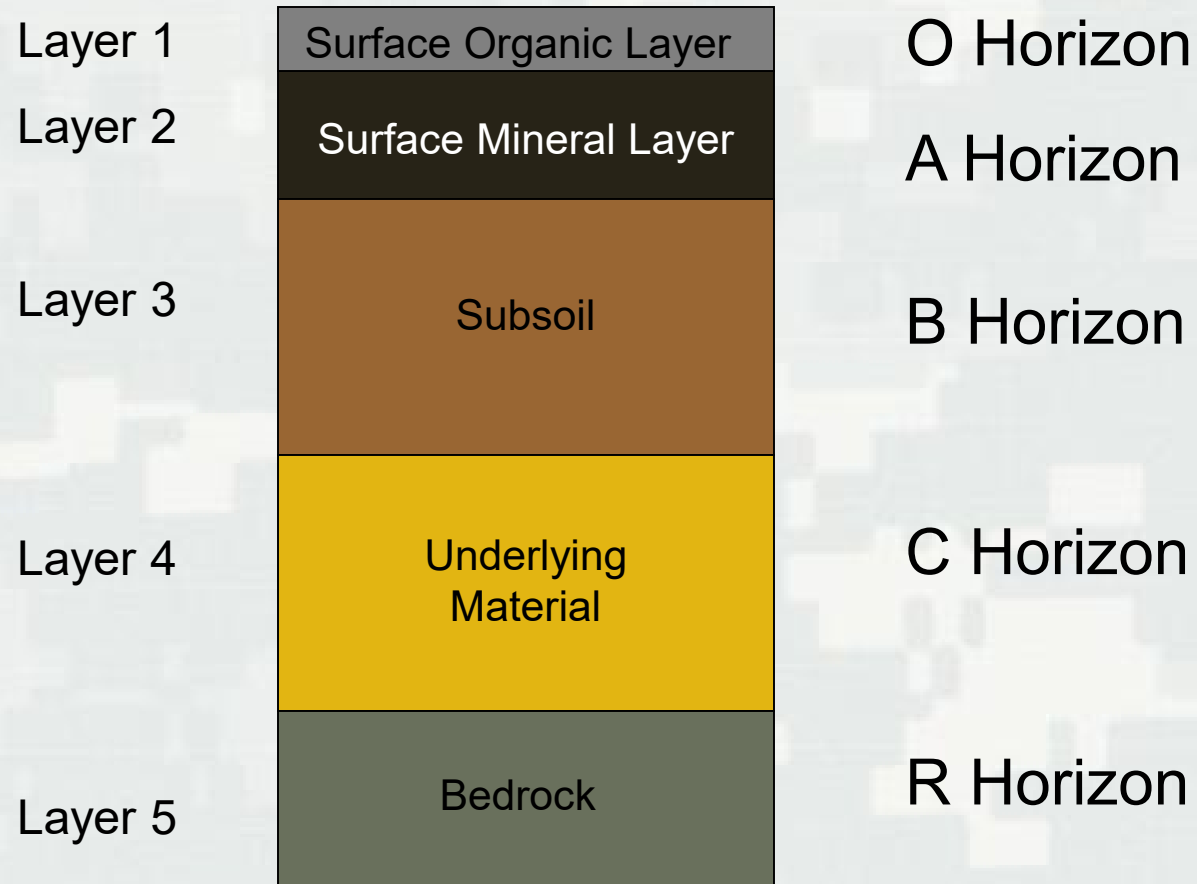


ERA	my	PERIOD	EPOCH
CENOZOIC	2	QUATERNARY	HOLOCENE PLEISTOCENE
	65	TERTIARY	PLIOCENE MIOCENE
		140	CRETACEOUS
MESOZOIC	210	JURASSIC	
	250	TRIASSIC	
	280	PERMIAN	PENNSYLVANIAN
PALEOZOIC	320	CARBONIFEROUS	MISSISSIPPIAN
	360	DEVONIAN	
	400	SILURIAN	
	440	ORDOVICIAN	
	500	CAMBRIAN	
	570	PRECAMBRIAN	

Soil Horizons

- Soils have different layers; these layers are called “horizons”. Major soil horizons are often discerned based upon soil color, organic matter content, and texture.
- A generalized soil profile might consist of the following horizons:
 - ▶ A horizon -- the surface layer in a mineral soil characterized by accumulation of organic matter and/or loss of materials (e.g., clays) to deeper layers.
 - ▶ B horizon -- the subsoil, characterized by accumulation of clays or other materials and greater structural development.
 - ▶ C horizon -- the underlying material, unconsolidated parent material little influenced by soil-forming processes.
 - ▶ R horizon – bedrock

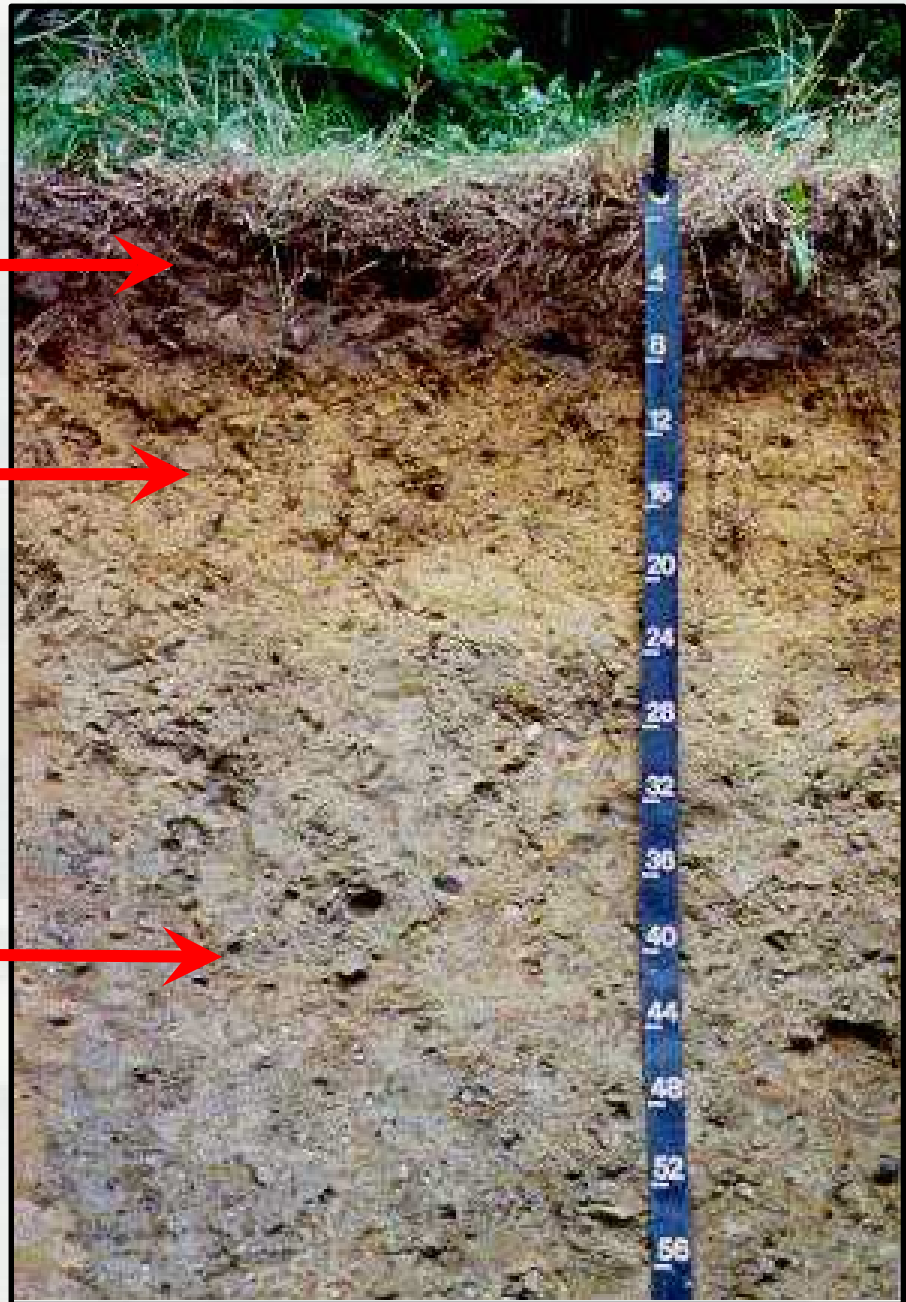
Major Horizon Designations



Layer 1 A Horizon →

Layer 2 B Horizon →

Layer 3 C Horizon →



Aspects of Soil Color

Munsell Soil Color System

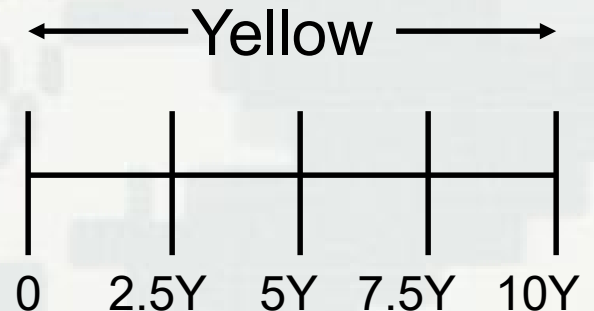
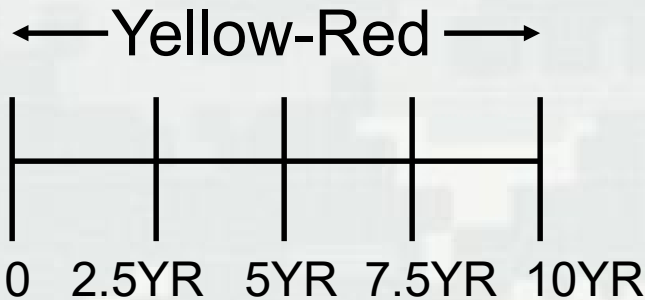
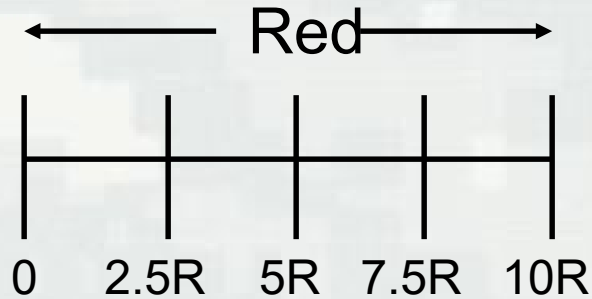
- Hue
- Value
- Chroma

10R 5/8



Hue

Spectral color in relation to red, yellow, green, blue, and purple.



Value

The **Lightness**
or **Darkness** of
Color

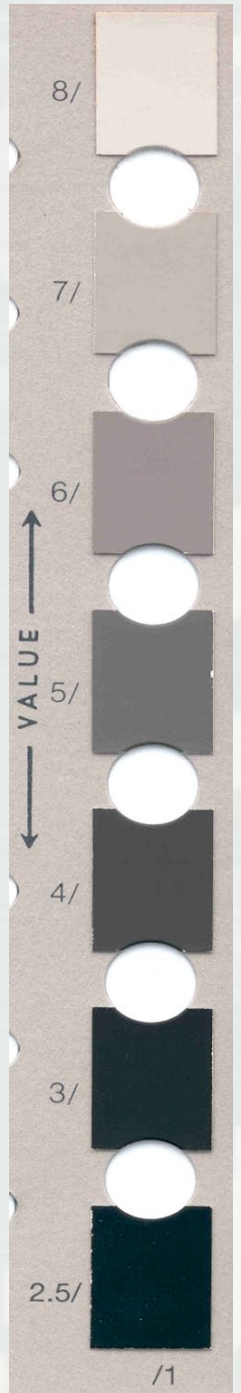
- 10/0 - Pure White



- 5/0 - “Gray”

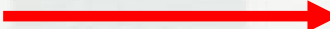


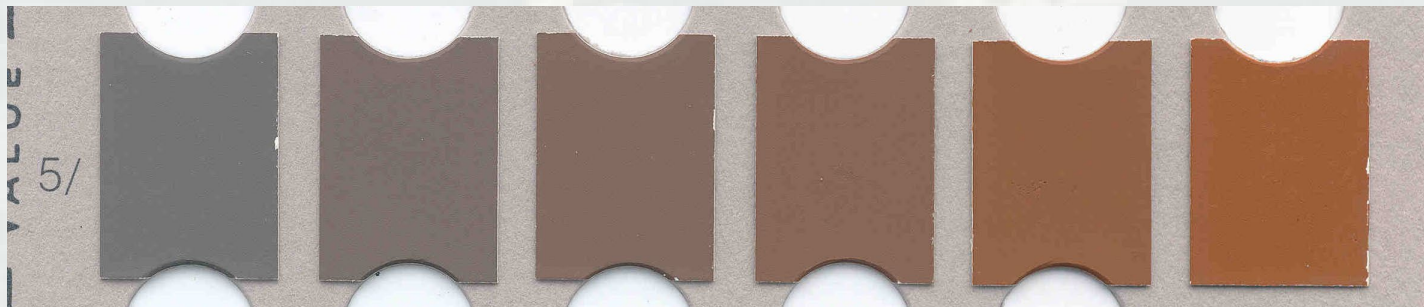
- 0/0 - Pure Black

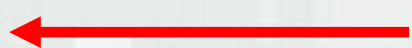


Chroma



Increasing strength of color 



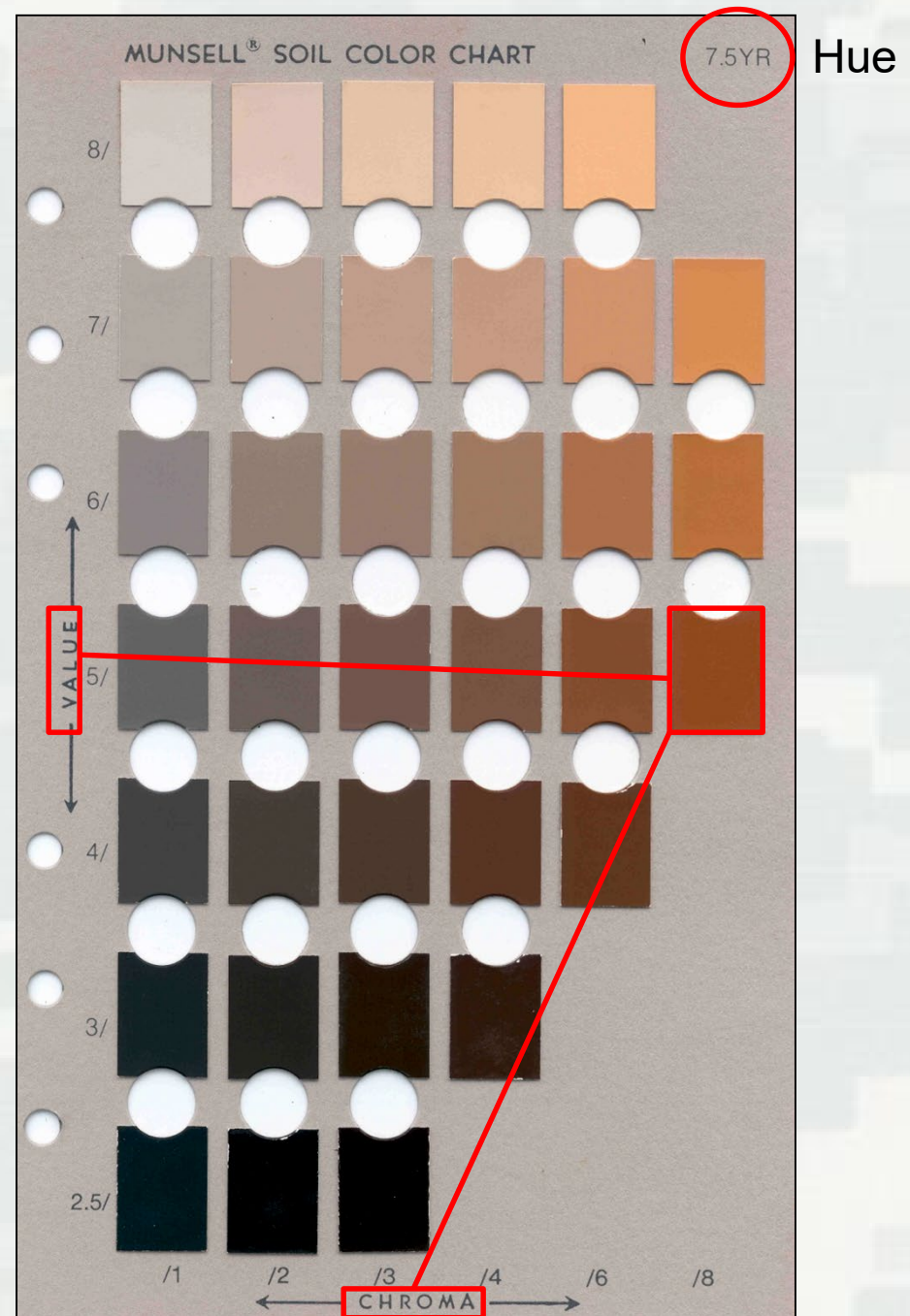
 Increasing grayness

Soil Color

Munsell Soil Color
Book, 7.5YR page

The highlighted color chip is:

7.5YR 5/8



Chroma

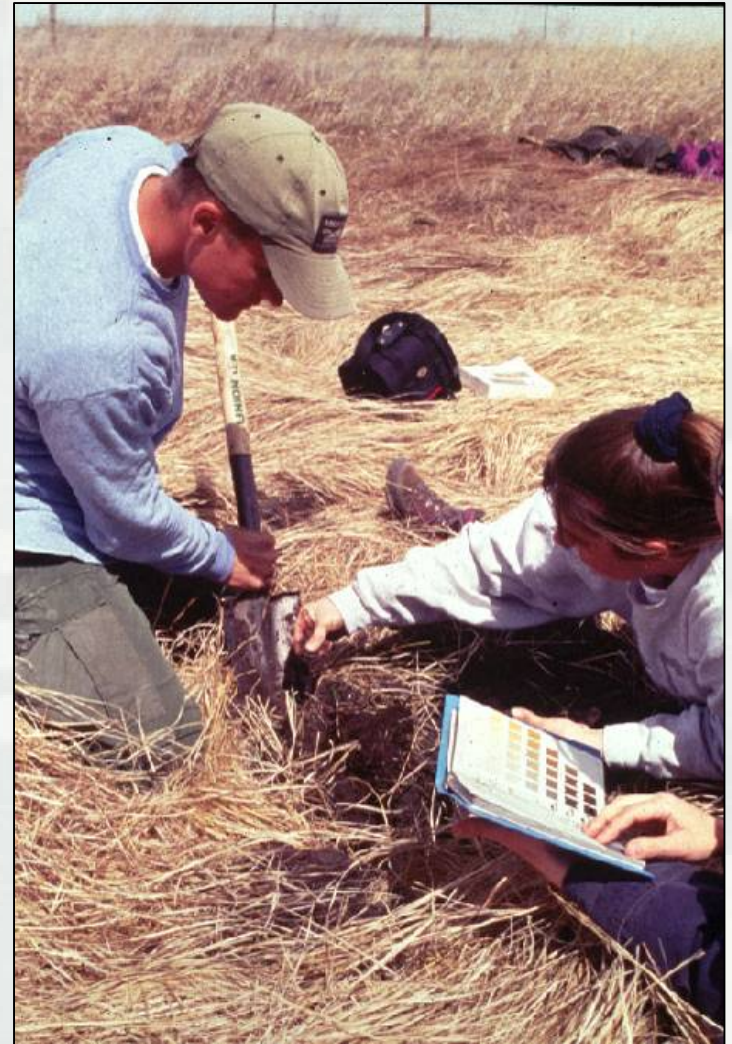
- Colors specified in the indicators do not list decimal points
- However, colors do occur between Munsell chips
- Do not round color to qualify as meeting an indicator



Reading Soil Colors

Optimum conditions for reading soil colors

- Natural light
- Clear, sunny day
- Midday
- Light at right angles
- Soil moist



Moist Colors

All color requirements (hue, value, and chroma) are for moist soils

- If dry, moisten to record color
- If wet, allow to dry to a moist state



Moist

Dry

Color Patterns in Soils

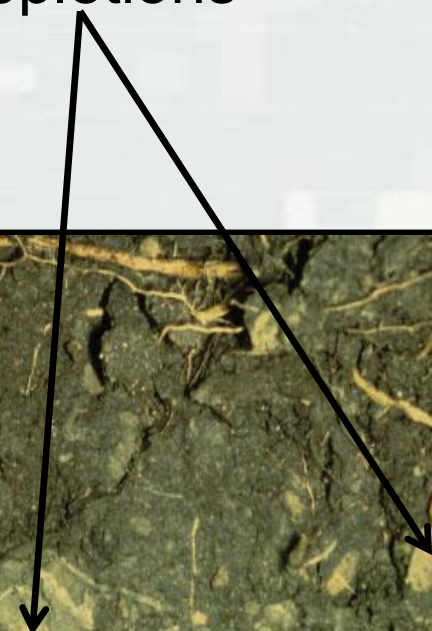
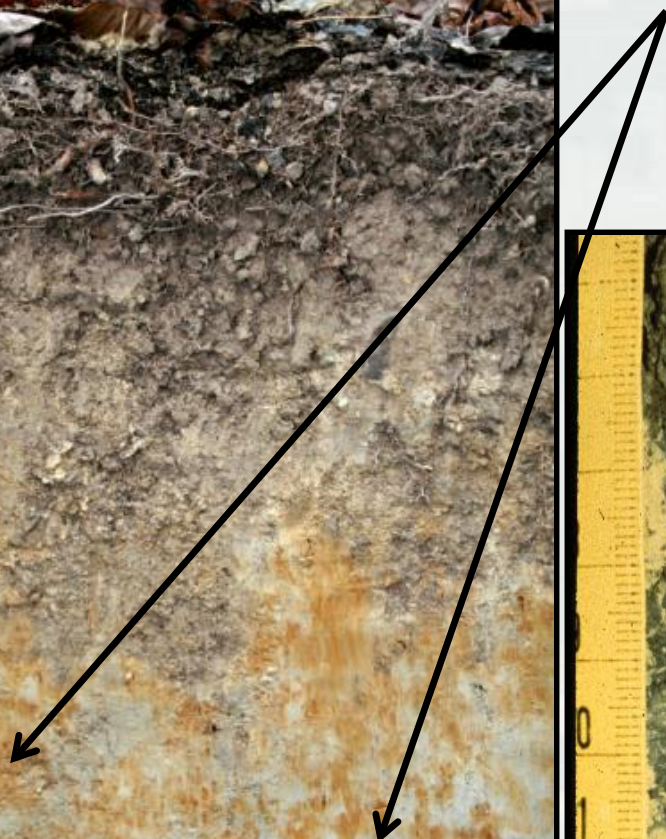
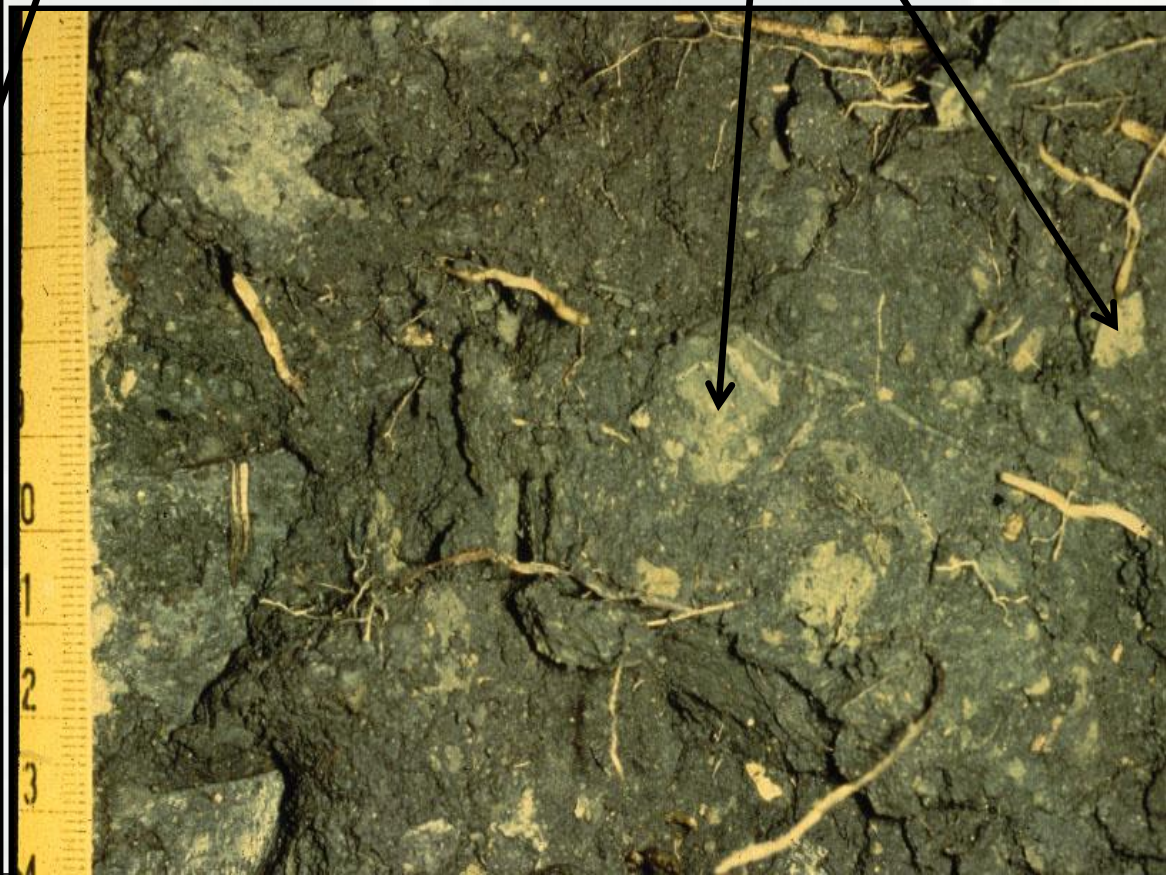
- Matrix (predominant) color
- Redoximorphic features
- Feature contrast and abundance





Redox Features Within a Soil Matrix

Concentrations -- Depletions



Contrast of Features

The degree of visual distinction between associated colors

- Faint – evident only on close examination
- Distinct – readily seen
- Prominent – contrast strongly

Upper Threshold for Faint

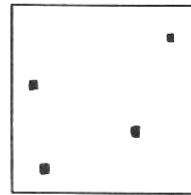
Delta Hue	Delta Value	Delta Chroma
0	≤ 2	≤ 1
1	≤ 1	≤ 1
2	0	0
Delta Hue	Delta Value	Delta Chroma
Any	≤ 3	≤ 2

Abundance of Features

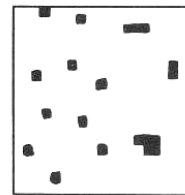
Different indicators require different abundance thresholds.

Best to record an estimated percentage.

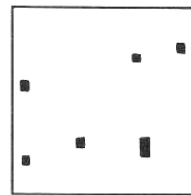
CHARTS FOR ESTIMATING PROPORTIONS OF MOTTLES AND COARSE FRAGMENTS



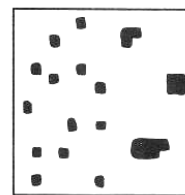
1%



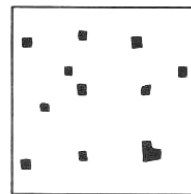
5%



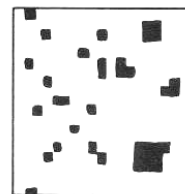
2%



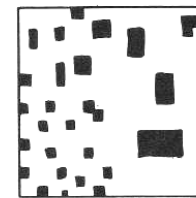
7%



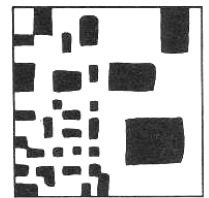
3%



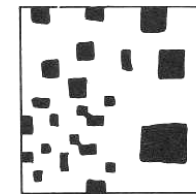
10%



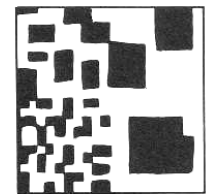
15%



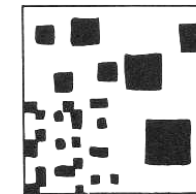
30%



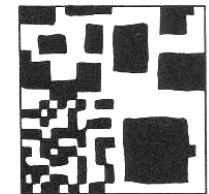
20%



40%



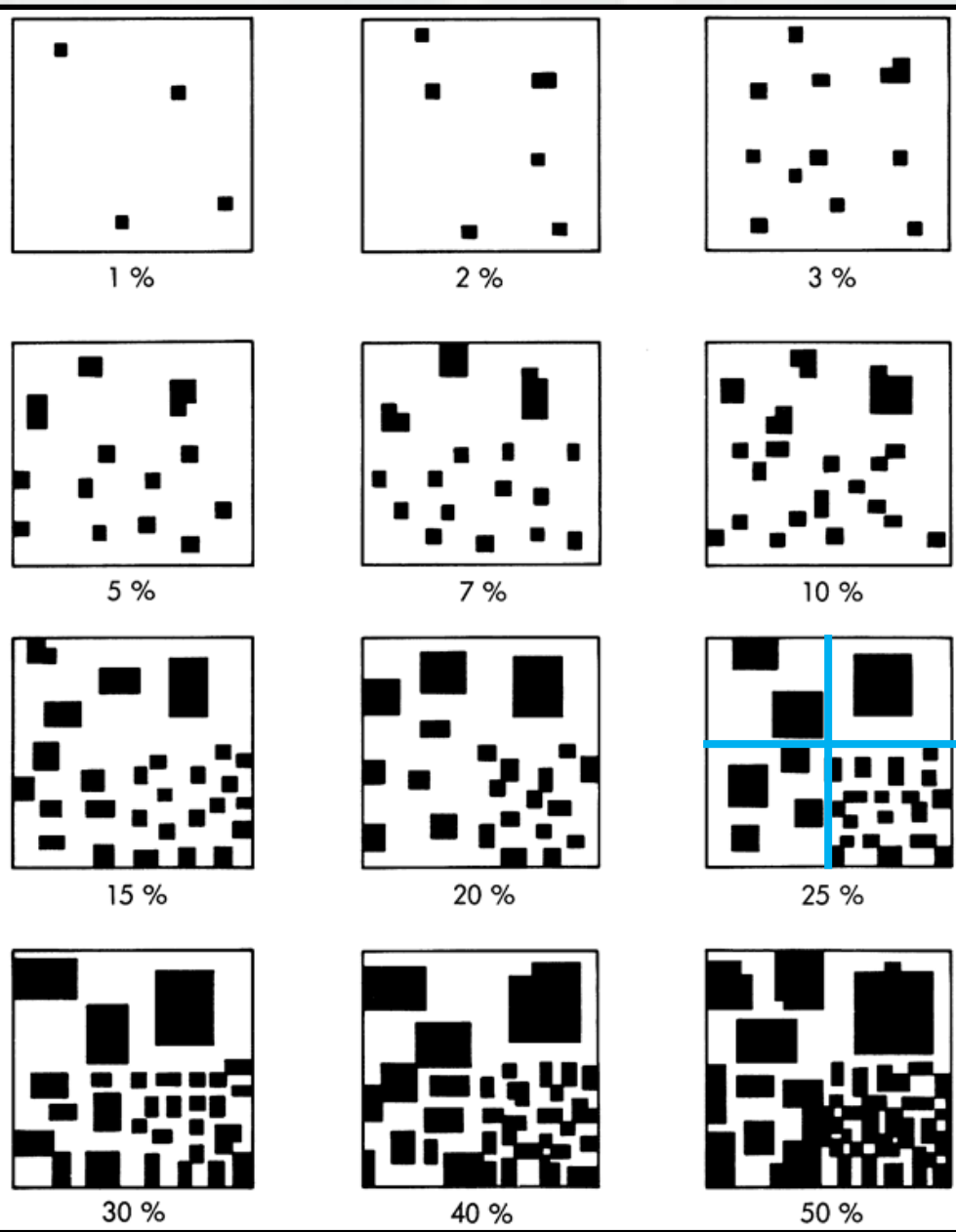
25%



50%

Each fourth of any one square has the same amount of black

Estimating Percent



Percentage of gray = 30%
Percentage of orange = 70%



Estimating Percent

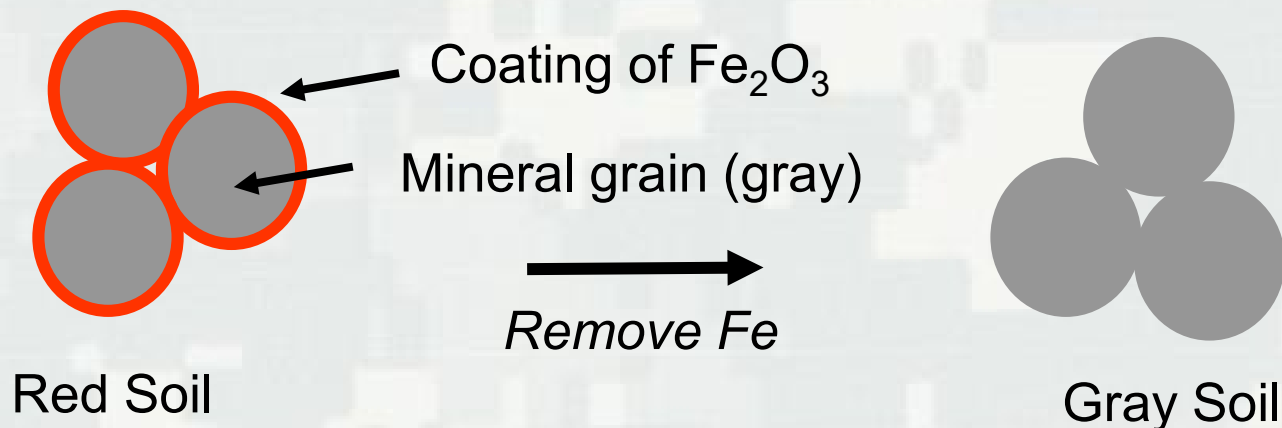
- Can all redox concentrations fit on one side?
- Can all redox concentrations fit in one corner?
- Breakdown further if necessary.



Oxidation/Reduction and Soil Color

[or, why concentrations and depletions look the way they do]

- In soil layers, Fe and Mn oxides give soils their characteristic brown, red, and yellow colors.
- When saturated and reduced, Fe and Mn are mobile and can be stripped from soil particles.
- This leaves the characteristic mineral grain color, usually a neutral gray.



Types of Redoximorphic Features

- Redox Concentrations
 - ▶ Masses
 - ▶ Pore Linings
 - ▶ Nodules and Concretions
- Redox Depletions
 - ▶ Depleted Matrix
- Reduced Matrix
- Coated sand grains

Redox Concentrations

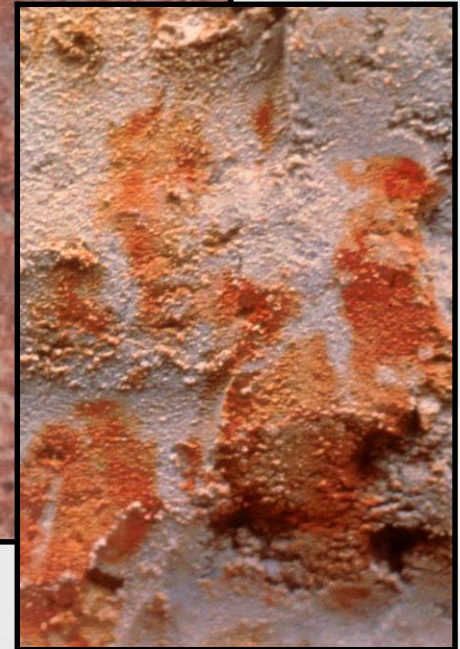
Bodies of apparent accumulation of Fe/Mn oxides

- Masses
- Pore linings
 - ▶ on ped faces
 - ▶ in root channels
- Nodules and concretions



Masses

- Soft bodies
- Frequently in the soil matrix
- Variable in shape



Pore Linings

- Coatings on a pore surface
- Impregnations of the matrix adjacent to the pore



Nodules and Concretions

- Firm to extremely firm bodies
- Often relict
- Contemporary ones should be irregular in shape
- Diffuse boundary
 - ▶ “halo” or “corona”



Fe/Mn nodules washed
from the soil

Relict vs Contemporary



Relict



Contemporary

Relict features are often firm to extremely firm and have abrupt boundaries with the soil matrix.

On left is a relict pore lining in deoxidized loess from eastern Nebraska. Commonly called pipe stems, these features formed at least 10,000 years ago (Pleistocene). On right, contemporary depletions along root channels and a contemporary Fe/Mn soft mass in lower left.

Age of Features

Redox features do not always indicate current hydrologic conditions

- Can be relicts of past climates
 - ▶ relict features can persist for thousands of years (e.g., Pleistocene).
 - ▶ relict features may have sharp edges and abrupt boundaries
 - ▶ relict nodules and concretions are often rounded
- Contemporary features should have diffuse boundaries and/or be associated with ped faces or root channels

Redox Depletions

Bodies of low chroma where Fe/Mn oxides have been stripped out

- Generally value ≥ 4
- Chroma ≤ 2
- Formerly called “gray mottles”



Depleted or Reduced Matrix

- The concept of a depleted matrix is that due to prolonged saturation, the soil has had significant iron removed, leaving the neutral base gray color as dominant. This is different from a reduced matrix.
- A reduced matrix occurs when soils have high value and low chroma *in situ*, but color changes when exposed to air. Reduced Fe^{2+} is present and oxidized to Fe^{3+} upon exposure to air. Specifics of color change requirements from Chapter 5 are discussed later.

Depleted Matrix

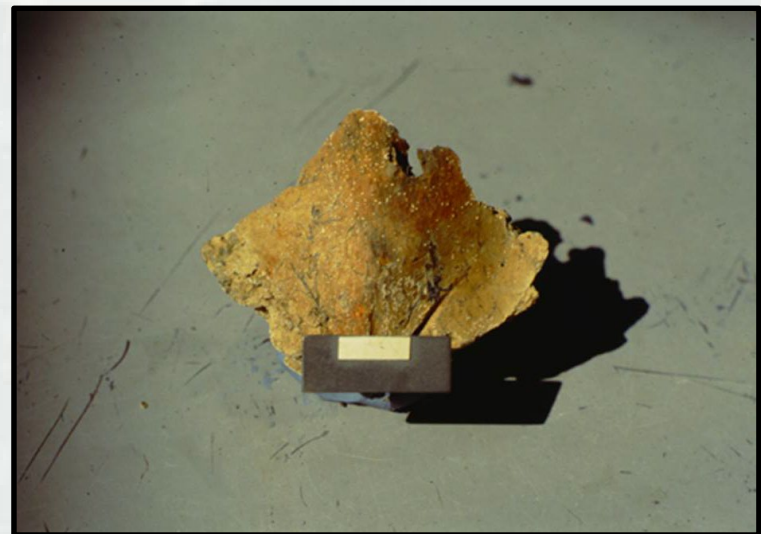
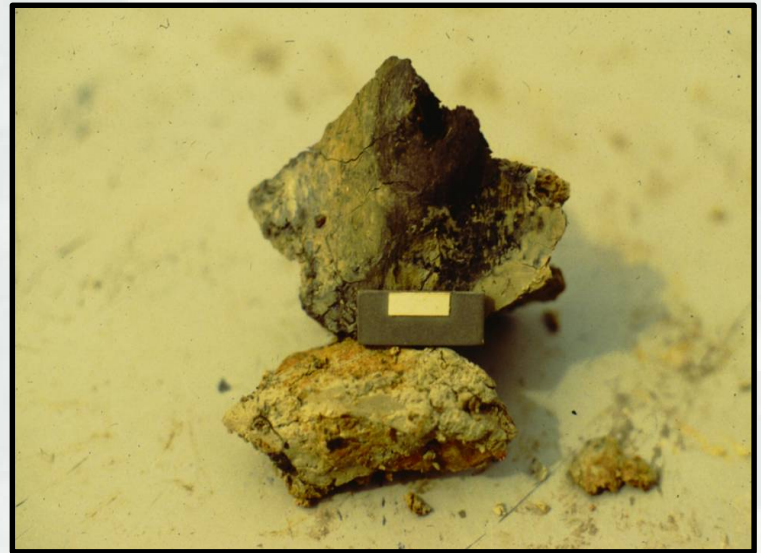
- Dominant color of the soil is “gray” due to removal (depletion) of iron
- Commonly used to identify hydric soils
 - ▶ Discussed more in the material on hydric soil indicators



Reduced Matrix

Soils have high value and low chroma *in situ*, but color changes when exposed to air

- Reduced Fe is present
- Fe^{2+} is oxidized to Fe^{3+} upon exposure



Redox Reactions

How do they help identify hydric soils?

<u>Element</u>	<u>Sign that Reduction Occurred</u>
Oxygen	Organic carbon accumulation
Nitrogen	None
Manganese	Redoximorphic Features
Iron	Redoximorphic Features
Sulfur	Rotten egg odor
Carbon	Methane

These reactions make wet soils look or smell distinctive allowing development of hydric soil indicators

Interpretation Problems

Not all wet soils develop redoximorphic features

- Low amounts of soluble organic carbon
- High pH
- Cold temperatures
- Low amounts of Fe
- Aerated groundwater

Rate of Feature Formation

- Soils in recently constructed wetlands develop visible redox depletions within a couple years if saturation occurs during the growing season
- Redox concentrations (pore linings) begin forming around aerated roots in as little as 7 days in flooded rice paddies, but may take 30 days or more before they start to become visible to the naked eye

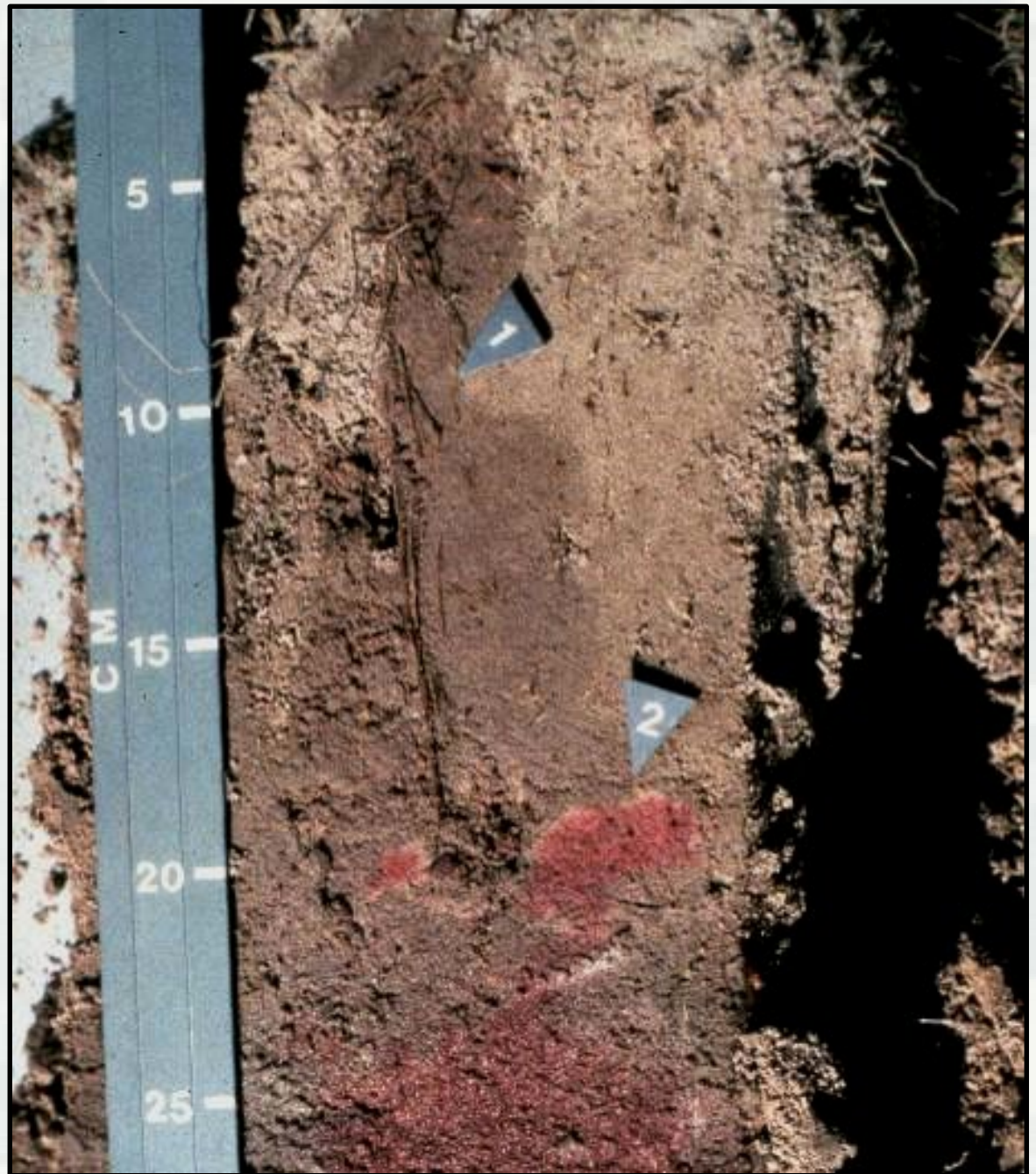
Alpha, Alpha-Dipyridyl

A dye used to test for the presence of reduced Fe

- Pink reaction to Fe^{2+}
- Dye is sensitive to light and heat
- Apply to freshly broken soil ped



Alpha, alpha-dipyridyl
applied to a
fresh slice of
soil



Alpha, Alpha-Dipyridyl Test Strips



Describing Redoximorphic Features

- Concentrations and Depletions
 - ▶ Describe type, color, abundance and location (i.e., along macropores or within matrix)
 - contrast can be obtained from color charts
 - automatically determined on automated data sheets (i.e., distinct/prominent)
- Reduced Matrix
 - ▶ Describe reduced matrix color, oxidized color, and time for color change to occur
- Alpha, Alpha-Dipyridyl
 - ▶ Describe % of soil that reacts and location

Two Categories of Soil Material

Differ in organic matter content:

- Organic material
- Mineral material

Organic Layers

- Consist of decomposed organic material
 - ▶ Peat (fibric) – least decomposed
 - ▶ Mucky peat (hemic)
 - ▶ Muck (sapric) – most decomposed

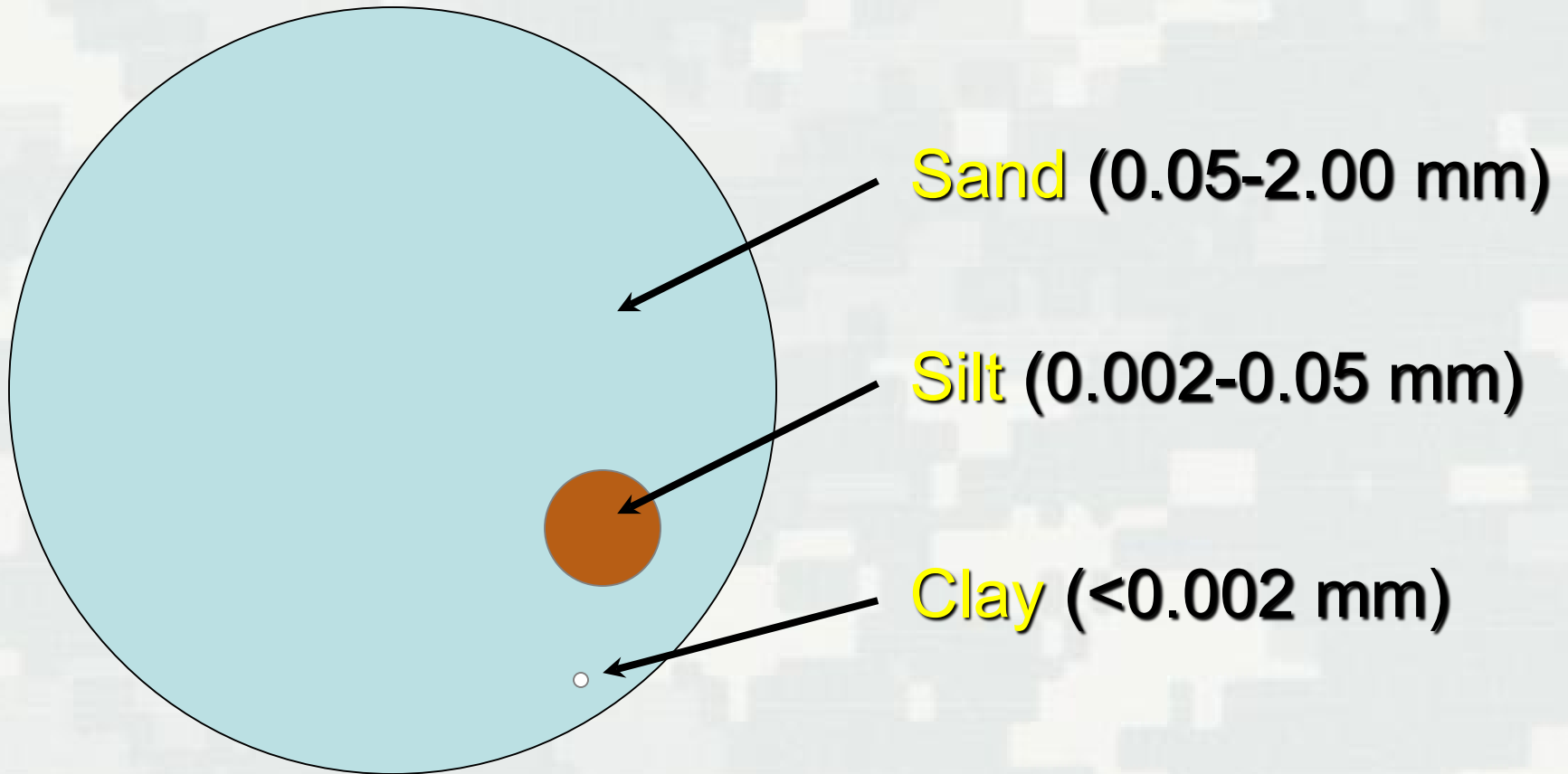


Mineral Layers

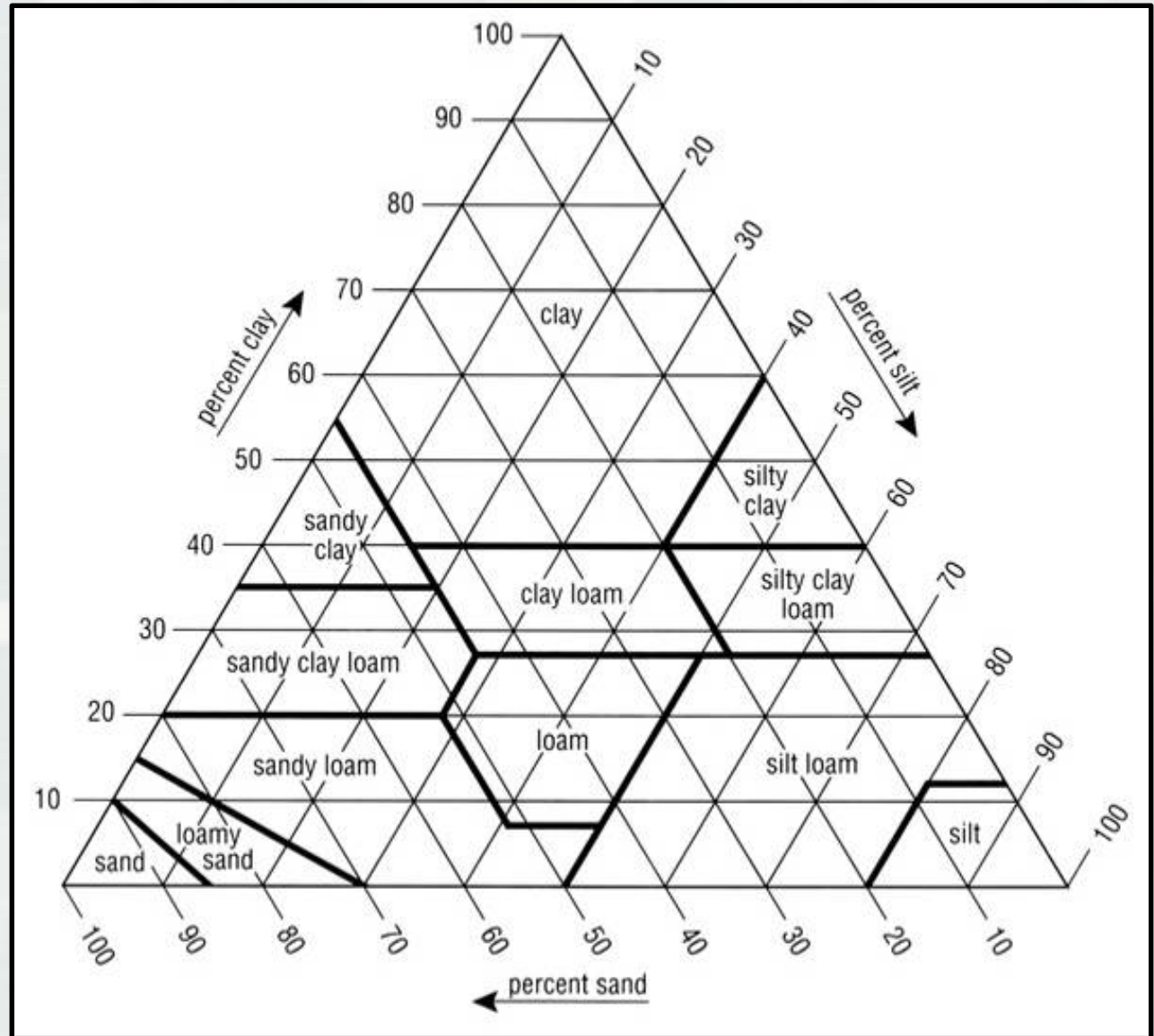
- Primarily sand, silt, and clay, with varying amounts of organic matter



Relative Sizes of Soil Particles

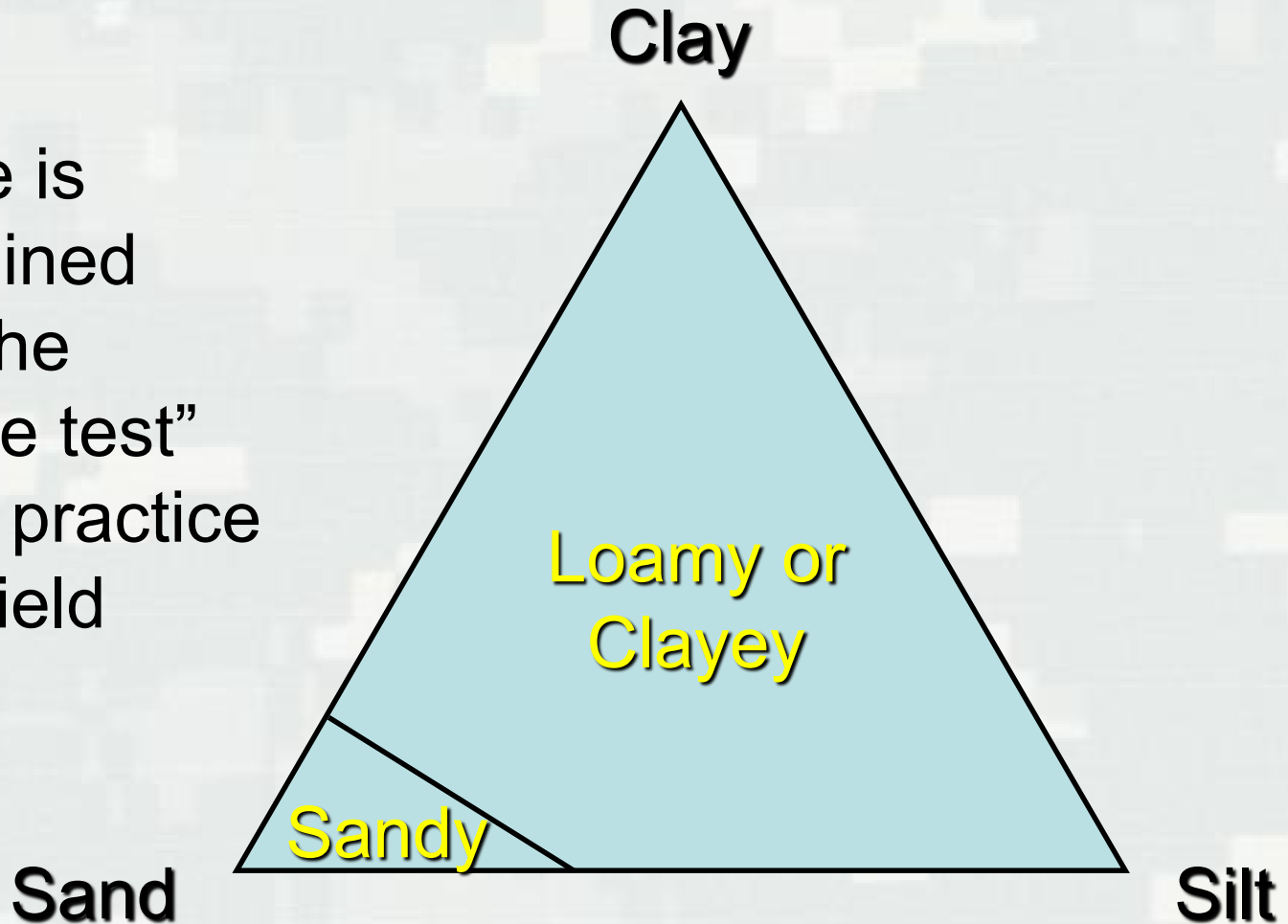


Soil Texture Triangle

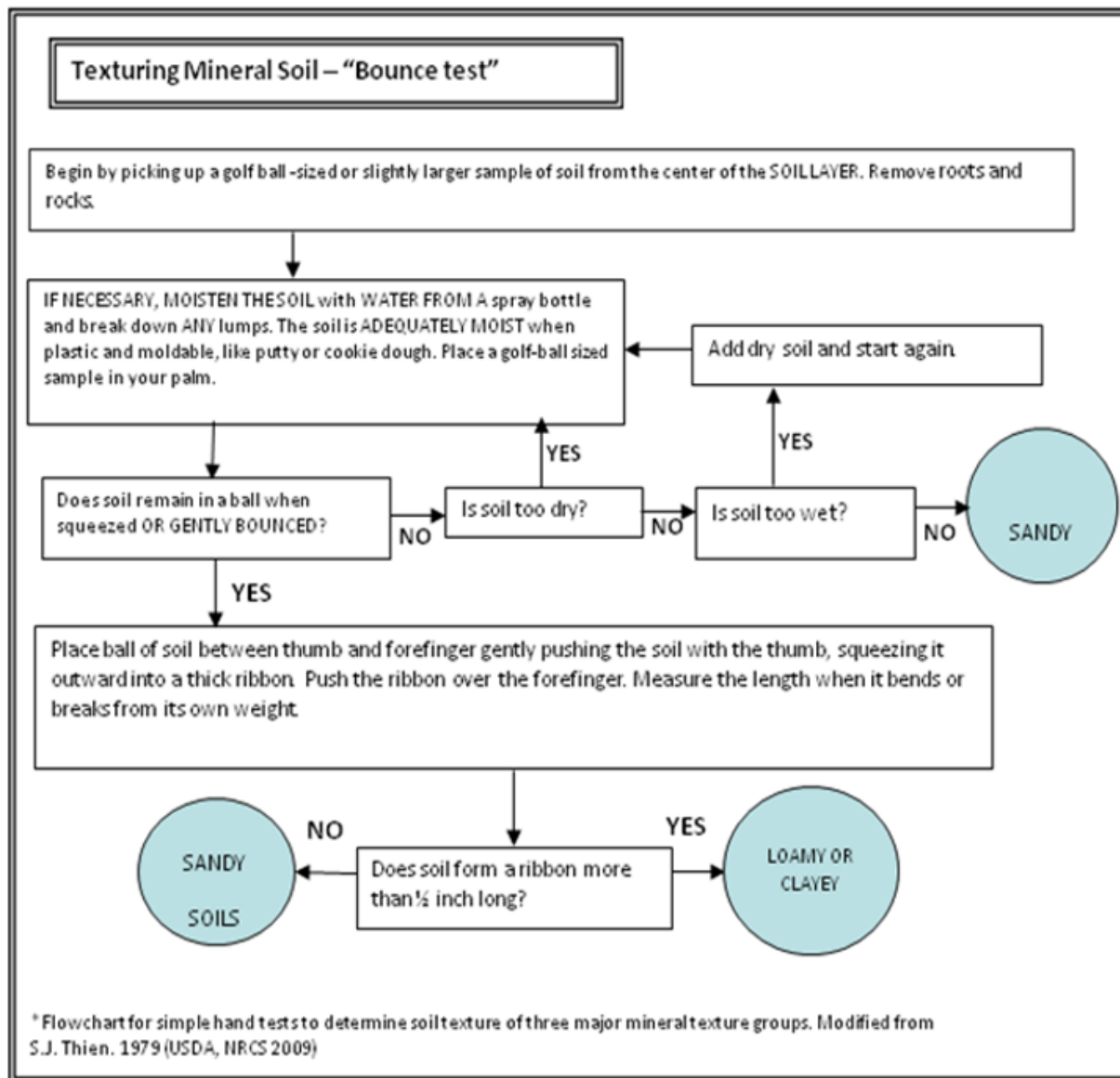


Simplified Soil Texture Triangle

Texture is determined using the “bounce test” we will practice in the field



Bounce Test



Where do soil observations begin?

The soil in this example begins:

- **beneath the knife blade**
- **below the fibric and hemic material**



Describe the Soil Profile

To select the appropriate indicator(s), it is critical to accurately describe the soil profile on the data sheet.

The soil profile is layered.

Each layer...

- ...has a thickness (and depth),
- ...has a color,
- ...has a texture, and
- ...may have other features.

